

High-resolution Atmospheric $p\text{CO}_2$ Reconstruction across the Paleogene Using Marine and Terrestrial $\delta^{13}\text{C}$ records

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The early Paleogene (63 to 47 Ma) is considered to have a greenhouse climate¹ with proxies suggesting atmospheric CO_2 levels ($p\text{CO}_2$) approximately $\sim 2\times$ pre-industrial levels. However, the proxy based $p\text{CO}_2$ reconstructions are limited and do not allow for assessment of changes in $p\text{CO}_2$ at million to sub-million year time scales. It has recently been recognized that changes in C_3 land plant carbon isotope fractionation can be used as a proxy for $p\text{CO}_2$ with quantifiable uncertainty². Here, we present a high-resolution $p\text{CO}_2$ reconstruction ($n = 597$) across the early Paleogene using published carbon isotope data from both terrestrial organic matter and marine carbonates. The minimum and maximum $p\text{CO}_2$ values reconstructed using this method are broad (i.e., $170 +60/-40$ ppmv to $2000 +4480/-1060$ ppmv) and reflective of the wide range of environments sampled. However, the large number of measurements allows for a robust estimate of average $p\text{CO}_2$ during this time interval ($\sim 400 +260/-120$ ppmv), and indicates brief (sub-million-year) excursions to very high $p\text{CO}_2$ during hyperthermal events (e.g., the PETM). By binning our high-resolution $p\text{CO}_2$ data at 1 million year intervals, we can compare our dataset to the other available $p\text{CO}_2$ proxies. Our result is broadly consistent with $p\text{CO}_2$ levels reconstructed using other proxies, with the exception of paleosol-based $p\text{CO}_2$ estimates spanning 53 to 50 Ma. At this timescale, no proxy suggests $p\text{CO}_2$ higher than 2000 ppmv, whereas the global surface ocean temperature is considered to be $>10^\circ\text{C}$ warmer than today. Recent climate modeling suggests that low atmospheric pressure during this time period could help reconcile the apparent disconnect between $p\text{CO}_2$ and temperature and contribute to the greenhouse climate³.

References

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3. Poulsen, C.J., Tabor, C., White, J.D., 2015. *Science* 348, 1238-1241.